

1 **Comparison of various GPS processing solutions towards an efficient validation of the**
2 **Hellenic vertical network:**

3 **The E.LE.V.A.T.I.ON project**

4

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7

8 **Abstract**

9 A research project for the validation of the Hellenic vertical network is currently in progress.
10 Two investigation areas in Central and Northern Greece have been chosen. The areas include
11 several benchmarks of the national trigonometric and leveling networks. Static Global
12 Positioning System (GPS) observations as well as classical terrestrial leveling are performed
13 to assess the internal accuracy of the two networks. Some numerical tests based on GPS and
14 leveling measurements are presented and the goals of the project are outlined. The strategies
15 followed in the processing of GPS data are presented with emphasis on their future use to the
16 project evolution. GPS observations have been processed using various commercial as well
17 as scientific software packages in order to examine the influence of the processing algorithms
18 to the final results. Significant differences between the results of the various software

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19 packages have been revealed, particularly in the case of challenging observation conditions.
20 Finally, comparisons of the estimated geoid heights at GPS benchmarks (BMs) to EGM2008
21 geoid information are performed as a first step towards the evaluation of the Hellenic vertical
22 network. These comparisons indicate that the two investigations areas are of different internal
23 accuracy namely 8.3 cm and 15.8 cm in terms of sd of the differences at Attica and
24 Thessaloniki test areas.

25

26 **1. Introduction**

27 The determination of the 3-D positions is feasible nowadays with particularly high accuracy
28 using modern Global Navigation Satellite System (GNSS) positioning techniques. In contrast,
29 the determination of vertical positions is much more demanding, mainly due to the inherent
30 connection between the vertical reference systems and the earth's gravity field. Height
31 information, reckoned from an equipotential surface, is of particular importance for a variety
32 of applications from coastal management to construction and monitoring of technical works
33 like highways, railways, metros and bridges.

34 A project for the validation and quality control of the Hellenic vertical network named
35 **E.L.E.V.A.T.I.O.N** (Evaluation of the HeLEnic Vertical network in the frAme of the
36 European sysTEms and control networks InterconnectiON – Application in the areas of
37 Attica and Thessaloniki) is currently in progress. Two investigation areas, one in Attica
38 (Central Greece) and another in Thessaloniki (Northern Greece) have been chosen. The areas
39 include several height benchmarks (BMs) of the national trigonometric and leveling
40 networks. Static GPS observations as well as classical spirit leveling in combination with
41 trigonometric leveling are performed to assess the internal accuracy of the two networks.
42 Numerical tests based on GPS and leveling measurements are presented and the goals of the
43 project are discussed. The control and re-evaluation of the Hellenic vertical network is the

44 main objective of the proposed project. Height information of high accuracy and reliability in
45 a common reference system is essential. Especially today, with the pan-European effort for
46 the establishment of a common European Vertical Network (Sacher et al., 2007), the
47 validation of the Hellenic network seems a prudent decision. In order to underline the
48 importance of reference system unification, it should be mentioned that the International
49 Association of Geodesy (IAG) established a Special Study Group (SSG) for the connection of
50 various reference systems in Europe. This SSG (EUREF – <http://www.euref.eu>) since 1989
51 has introduced the European Terrestrial Reference System of 1989 (ETRS89). The
52 connection of the Hellenic 3-D network with ETRS89 has been established through the
53 Hellenic Positioning System (HEPOS). HEPOS is a nation-wide Real Time Kinematic (RTK)
54 network based on 98 reference stations established for the modernization of the geodetic
55 infrastructure of Greece (Gianniou, 2008). During the next years the connection of the
56 vertical datum with Europe has to be done; this is also a European Community directive
57 under the name “INSPIRE”. Before the connection, the validation of the vertical network has
58 to be carried out.

59 The first-order vertical control network of Greece was established and measured by the
60 Hellenic Army Geographic Service from 1963 to 1986 (Miloni – Kotroyanni, 1989).
61 Approximately 11000 km of traverses and 11000 vertical control benchmarks are the
62 characteristics of Greek vertical network. The tide gauge in Piraeus harbor is the fundamental
63 point of the network. On the other hand, the first order Hellenic trigonometric network has
64 some height information, due to some trigonometric leveling lines. This vertical information
65 has not been validated since its creation. The validation of the vertical reference network
66 before the establishment of the European interconnection is thus essential.

67 The European committee for the continental control networks works under the auspices of the
68 European Council on the measurement and establishment of both a horizontal as well as a

69 vertical European reference system. A Vertical System is characterized by its Datum (point of
70 reference) and the type of height used. The Datum point is estimated by the Mean Sea Level
71 (MSL) in the area, as determined by tide gauge measurements. In Europe tide gauges exist in
72 various regions: in the Baltic, in the North Sea, in the Mediterranean, in the Black Sea and in
73 the Atlantic Ocean. Level differences between various tide gauges can reach several
74 centimeters. In addition, national vertical datum points are based on historical facts and not
75 always referenced to the MSL, e.g., the zero-point of the Amsterdam tide gauge is defined as
76 the mean high tide in the year 1684.

77 Another issue is the use of various types of heights around Europe. Thus, orthometric heights
78 are used in Belgium, Denmark, Italy, Greece, etc. and normal heights are used in France,
79 Germany, Sweden and the Eastern European countries. In 1945 the integration of national
80 systems started, while the establishment of a common system around Europe was divided to
81 various solutions in Western and Eastern Europe due to political reasons.

82 Greece, in particular, has not been connected yet with any of the unified vertical reference
83 systems. As a consequence, difficulties arise in planning and executing cross-border works
84 like roads, railways and pipeline constructions. A prerequisite for the Hellenic vertical datum
85 connection is its evaluation. The validation of the height data must be based on the
86 interpretation of the inner accuracy of the solution and the external control using independent
87 data.

88 The first stage of the ELEVATION project is dedicated to the compilation/validation of
89 existing data and the collection of new observations. These observations were collected
90 during the first stage of the project (August-October, 2012) and referred to the update and
91 enrichment of the existing GPS and leveling database. GPS observations near leveling
92 benchmarks (reperes) of the Greek vertical network as well as on trigonometric pillars were
93 collected. The connection between various benchmarks using classical spirit or trigonometric

94 leveling with simultaneous reciprocal-observations is also part of the first stage of the project.
95 The second stage of the project is based on the data processing. GPS observations have been
96 processed using various commercial as well as scientific software packages in order to
97 examine the influence of the processing algorithms to the final result.

98

99 **2. Theoretical background on heights**

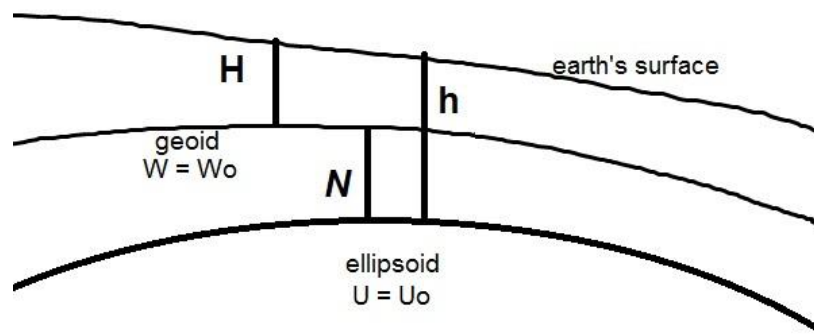
100 The need to separate horizontal and vertical positions stems from the different accuracy
101 provided by terrestrial observations. Horizontal directions are measured with increased
102 accuracy compared to the vertical ones. This is due to the atmospheric refraction effect. The
103 abovementioned fact introduces greater uncertainty to vertical positioning. This is why
104 classical geodetic observations are divided into horizontal directions and distances for
105 horizontal positioning and spirit leveling measurements for vertical positioning (Torge,
106 2001).

107 Height data are referenced to suitable level surfaces, which represent characteristic elements
108 of the observation environment. Heights are connected with human activity and thus their
109 link with physical characteristics is necessary. A characteristic surface is the Mean Sea Level
110 (MSL). This surface represents the traditional connection of all human activities with the
111 natural environment. Practically, it is common knowledge that the MSL is a zero-height
112 surface. Theoretically speaking, MSL in a global scale constitutes a balance surface of waters
113 and, excluding the presence of the quasi-stationary Dynamic Ocean Topography, represents
114 an equipotential surface of Earth's gravity field. In this manner, the concept of geoid as a
115 height reference surface is introduced. The geoid is an equipotential surface of the Earth's
116 gravity field that to a first approximation coincides with the MSL in global scale, provided
117 that the effects of tides and ocean currents are removed. In a well-defined national vertical
118 control network, heights are referenced in a datum point of zero altitude. Usually, the zero-

119 height point is defined by local MSL observations from tide gauge records. In reality, the sea-
120 level change is measured from a conventionally selected level, which is considered constant:
121 the tide gauge zero.

122 Another reference surface used is the ellipsoid of revolution. The ellipsoid is not a physical
123 surface and is used only as a model of the Earth's surface for horizontal positioning, due to
124 the simplicity of its mathematical description. Data from geodetic satellite missions can be
125 referenced to an ellipsoid of revolution. The data of such missions will be used for the
126 validation of the current vertical network. The main height reference surfaces used in this
127 work are depicted in Figure 1 and analytically described in the methodology section.

128



129

130 **Figure 1:** Height reference surfaces

131

132 A point in space can be identified using three coordinates: the latitude, the longitude and the
133 height. The horizontal coordinates are referenced to the surface of a model ellipsoid of
134 revolution which is a geometrical-mathematical surface related to the MSL in local or global
135 scale. The height of a point P can be referenced along the vertical on the ellipsoid and is
136 called ellipsoidal height h_p .

137 However, in geosciences the altitude of a point must be referenced to the MSL, or more
138 precisely to the vertical reference system. As it is known, the ellipsoidal model does not

139 coincide with the MSL but has a deviation from -100 to 100 m, globally. The dependence of
 140 the vertical reference system from the gravity field seems obvious, since an equipotential
 141 surface of this field is the first approximation of the MSL at a global scale.

142 The earth gravitational potential is the potential of the attracting masses including the
 143 atmosphere and can be expressed in spherical harmonic expansion (Hofmann-Wellenhof and
 144 Moritz, 2005):

$$145 \quad V_e(r, \theta, \lambda) = \frac{GM}{R} \sum_{n=0}^{\infty} \sum_{m=-n}^n \left(\frac{R}{r}\right)^{n+1} C_{nm} \bar{Y}_{nm}(\theta, \lambda), \quad (1)$$

146 where (r, θ, λ) are the spherical coordinates of the computation point, GM is the product of
 147 Newton's gravitation constant and Earth's mass (including the atmosphere), R is the mean
 148 radius of the Earth, C_{nm} a constant coefficient of degree n and order m , \bar{Y}_{nm} are the fully
 149 normalized spherical harmonic functions:

$$150 \quad \bar{Y}_{nm}(\theta, \lambda) = \begin{cases} \bar{P}_{nm}(\cos \theta) \cos |m| \lambda, & m \geq 0 \\ \bar{P}_{n|m|}(\cos \theta) \sin |m| \lambda, & m < 0 \end{cases}, \quad (2)$$

151 and \bar{P}_{nm} are the fully normalized Legendre functions of the first kind. The above expression
 152 is the solution of the boundary value problem for the potential and it is valid for any point
 153 outside the Earth masses, in which space Laplace's equation is applied (Martinec, 1998).

154 Due to Earth's rotation, the gravimeters measure additionally a centrifugal acceleration
 155 which leads to the centrifugal potential:

$$156 \quad \Phi(r, \theta) = \frac{1}{2} \omega_e^2 r^2 \sin^2 \theta, \quad (3)$$

157 where ω_e is the angular velocity of Earth's rotation. Therefore the gravity potential can be
 158 expressed as:

$$159 \quad W(r, \theta, \lambda) = V(r, \theta, \lambda) + \Phi(r, \theta). \quad (4)$$

160 The connection between geometrical and physical characteristics is established following the
161 equation:

$$162 \quad \mathbf{g} = \nabla W \quad (5)$$

163 and the magnitude of the gravity vector:

$$164 \quad |\mathbf{g}| = g = -\frac{dW}{dn}, \quad (6)$$

165 where dn is the differential length along the plumb-line.

166 As it is already mentioned, a height reference surface must be related with the physical
167 environment through an equipotential surface of Earth's gravity field, a surface of constant
168 value of W . Especially, the surface $W=W_0$, which is approximated by the MSL is known as
169 the geoid. Therefore, the height "above MSL" is defined precisely as the height "above the
170 geoid". Let $P_0^{(j)}$ a point near a tide gauge, with a gravity potential $W_0^{(j)}$. There exist three
171 different kinds of heights depending on the potential definition at the point of interest. This
172 potential difference is known as geopotential number:

$$173 \quad C_P^{(j)} = W_0^{(j)} - W_P, \quad (7)$$

174 where W_P is the gravity potential of point P on Earth's surface. The geopotential number is
175 a unique characteristic of the space domain and using a scale factor of normal gravity γ_0 can
176 be expressed as height coordinate:

$$177 \quad H_P^{dyn(j)} = \frac{C_P^{(j)}}{\gamma_0}, \quad (8)$$

178 which represents the dynamic height of point P (related with the local vertical datum j).

179 The normal gravity scale factor is usually taken equal with the magnitude of normal gravity
180 computed at a mean latitude ($\gamma_0 = 9.806199203 \text{ m/s}^2$). It is noted that the dynamic height is
181 expressed in length units and can be used as a height. Nevertheless, it does not provide any

182 geometrical information: it is just a physical quantity – the potential related to the geoid
183 surface.

184 Seeking for a geometrical definition the integration of equation (8) is performed:

$$185 \quad W_P = W_0^{(j)} - \int_{P_0^{(j)}}^P g dn. \quad (9)$$

186 Using the geopotential number at point P it becomes:

$$187 \quad C_P^{(j)} = \int_{P_0^{(j)}}^P g dn. \quad (10)$$

188 Equation (10) shows the relation between the geopotential numbers, gravity and
189 measurements of vertical difference between equipotential surfaces along the plumbline. The
190 vertical differences (in length units) are observed using classical spirit leveling. Taking the
191 measured track always perpendicular with the equipotential surfaces (plumb-line), the
192 geopotential number can be computed as:

$$193 \quad C_P^{(j)} = \int_{\bar{P}^{(j)}}^P g dH, \quad (11)$$

194 where dH is the unit-length along plumb-line and $\bar{P}^{(j)}$ is located at the intersection of the
195 plumb-line with the geoid surface. Solving the above equation along the vertical length, $H_P^{(j)}$,
196 called orthometric height, one can write:

$$197 \quad H_P^{(j)} = \frac{C_P^{(j)}}{\bar{g}_P^{(j)}}, \quad (12)$$

198 where

$$199 \quad \bar{g}_P^{(j)} = \frac{1}{H_P^{(j)}} \int_{\bar{P}^{(j)}}^P g dH \quad (13)$$

200 is the average gravity along the plumb-line. In this specific case, a density model for the
201 masses inside the Earth is needed. This fact dictates the direct dependence of orthometric
202 height accuracy with the accuracy of the density model used.

203 The relation between ellipsoidal heights, measured from GNSS and orthometric heights is:

$$204 \quad h = H + N, \quad (14)$$

205 where h is the ellipsoid height along the vertical on the model surface, H is the orthometric
206 height from the geoid surface, measured along the plumb-line and N is the geoid undulation
207 (distance from the geoid to the ellipsoid) along the vertical on the ellipsoid. According to the
208 definition, the orthometric height is independent of the ellipsoid model used. However, the
209 geoid undulation is based on the ellipsoid choice because it is expressed as the difference
210 from a specific model. Geoid heights can be derived using local gravity information in
211 combination with global features provided by a geopotential model. The most recent global
212 geopotential model calculated from a special spectral combination of terrestrial and satellite
213 data is EGM2008 (Pavlis et al., 2012).

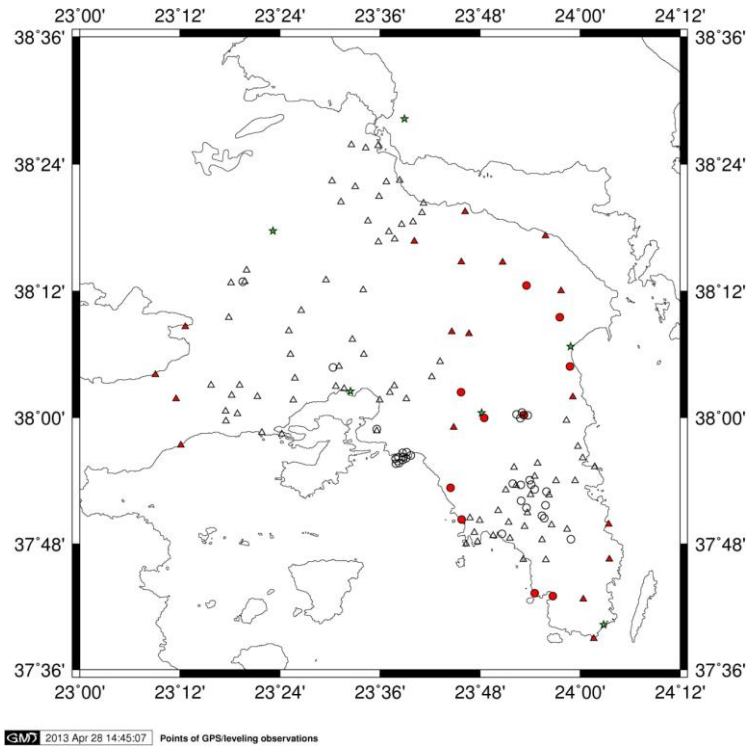
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215 **3. Data collection and analysis**

216 **3.1. GPS measurements**

217 The main purpose of the GPS measurements was the determination of the ellipsoidal heights
218 of the trigonometric and leveling BMs. As known, leveling BMs are often established on
219 vertical elements like walls or columns and, thus, they are not adequate for GPS
220 measurements. In such cases, we established new points offering good satellite visibility on
221 sites as close as possible to the original BMs (distances up to 200 m). These newly
222 established points were connected to the original BMs by means of double-run spirit leveling.
223 Figures 2 and 3 depict the location of the trigonometric and leveling benchmarks in Attica
224 and Thessaloniki, respectively.

225



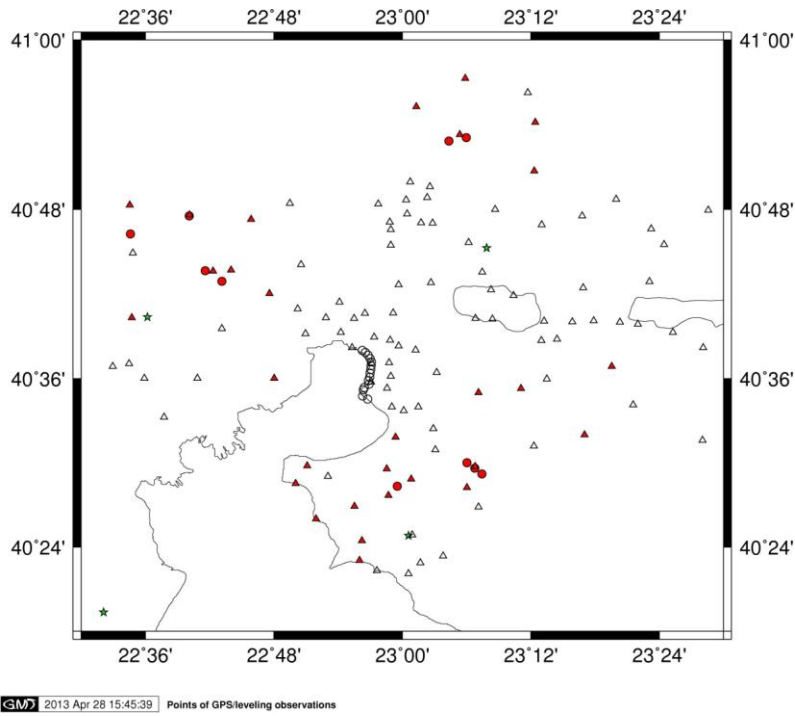
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227 **Figure 2:** Points in Attica region. Star: HEPOS stations, Circle: previously measured height

228 BMs, Solid Circle: newly measured BMs, Triangle: previously measured trigonometric BMs,

229 Solid Triangle: newly measured trigonometric BMs.

230



231

232 **Figure 3:** Points in Thessaloniki region. Star: HEPOS stations, Circle: previously measured
233 BMs, Solid Circle: newly measured BMs, Triangle: previously measured trigonometric BMs,
234 Solid Triangle: newly measured trigonometric BMs.

235

236 In order to ensure high accuracy in the determination of the ellipsoidal heights of the BMs,
237 the GPS measurements have been designed carefully. A key parameter for this work was the
238 selection of an adequate geodetic reference frame. The latest International Terrestrial
239 Reference Frame (ITRF2008) would be the best choice, as it ensures the highest possible
240 accuracy. However, this solution would require the connection of the BMs to permanent
241 reference stations with well known ITRF coordinates, i.e. IGS and/or EUREF/EPN stations.
242 As there are no IGS stations in Greece (except two proposed stations in Athens and Chania),
243 the length of the baselines to the closest IGS stations would be of the order of hundreds of
244 kilometers imposing observation times of at least 24-48 hours, which was improper for our
245 project. Regarding the EPN stations, NOA1 in Athens and AUT1 in Thessaloniki are situated
246 within the two project areas at distances up to 60 km away from the BMs. However, we
247 wanted to have every point connected to at least two stations, which would lead to baseline
248 lengths of the order of hundreds of kilometers. So, instead of using EPN stations, we
249 preferred to use stations of the Hellenic Positioning System HEPOS (Gianniou, 2008). The
250 system consists of a dense network of stations, offering the possibility to connect each BM to
251 two stations, while keeping short baseline lengths. The baselines measured in the area of
252 Attica and Thessaloniki are shown in Figures 4 and 5, respectively. The corresponding mean
253 baseline length for each area was 20.9 km and 23.4 km, respectively. The maximum baseline
254 length was 44 km and only four vectors among a total number of 134 baselines exceed 40
255 km. Given the aforementioned baseline length the rapid-static method could have been used.
256 However, in order to increase the accuracy of the results, the static method has been used

287 trigonometric leveling. More on the techniques used for the leveling observations as well as
 288 results and comparisons from the evaluation procedure can be found in Anastasiou et al.
 289 (2012).

290

291 **4. Data processing and results**

292 **4.1. GPS data processing schema**

293 Due to the challenging for GPS measurement environment at some BMs (foliages, obstacles,
 294 electromagnetic interferences), difficulties in the data processing had been expected. In order
 295 to have a better control on the quality of the results, it was decided to perform independent
 296 computations using five different software packages available at the three Institutions
 297 participating in the research project. In that manner it would be possible to make an extended
 298 comparison of the used software packages. Table 2 summarizes the programs used and their
 299 characteristics.

300

301 **Table 2:** Software packages used for the GPS data processing.

Software	Version	Release year	Manufacturer	Abbreviation in paper
Bernese	5.0	2007	University of Bern	BERN
Geomax Geo Office	2.0	2009	Geomax	GGO
GrafNet	8.40	2012	Novatel (XYZ solutions)	GN
Trimble Business Center	1.12	2007	Trimble	TBC
Topcon Tools	7.5.1	2010	Topcon	TT

302

303 For the processing with Bernese, IGS precise orbits have been used. For the processing with
304 the commercial software packages we used broadcast orbits. The error in the baseline length
305 introduced by the orbital error can be approximated by the formula (Teunissen and
306 Kleusberg, 1998):

$$307 \quad \frac{db}{b} = \frac{dr}{r} \quad (15)$$

308 where db/b is the relative baseline error and dr/r the relative orbital error. Given that the
309 maximum baseline length was 44 km and assuming an orbital error of 2 m, it comes out that
310 the maximum error in the baseline length due to the orbital error did not exceed 4 mm, which
311 is fully sufficient for our purposes.

312 With Bernese the processing parameters described in the CODE Analysis strategy
313 (ftp://ftp.unibe.ch/aiub/CODE/0000_CODE.ACN) were used. For fixing the ambiguities, the
314 SIGMA algorithm (Dach et al., 2007) has been used together with the L1/L2 method (for
315 baselines up to 20 km) and the widelane/narrowlane method (for longer baselines). With the
316 commercial software packages the default processing parameters of each software have been
317 used.

318

319 **4.2. GPS data processing results**

320 During the baseline processing we encountered certain difficulties due to the aforementioned
321 unfavorable satellite signal reception at some BMs. For several baselines the initial
322 processing (i.e. using all observed satellites) yielded fixed solutions but with poor statistics
323 (flagged fixed), whereas for a limited number of baselines the initial result was a float
324 solution. In order to improve the results we reprocessed these baselines after rejecting
325 observations with large residuals. In this way, most of the flags were removed and most of
326 the float solutions became fixed. This procedure was followed with the commercial
327 programs, which are suitable for such kind of interventions. In contrast, such intrusions are

328 quite complex in Bernese, so no similar attempts were made with this software. Table 3
 329 summarizes the initial and final results obtained from each processing software.

330

331 **Table 3:** Baseline solution results.

Initial Solution	Final Solution	Comments	# of baselines per software				
			BERN	GGO	GN	TBC	TT
Float	Float	Solution not used	6 *	1	-	-	-
Float	Fixed	Fixed obtained after deactivating satellites	- *	3	-	6	7
Fixed Flagged	Fixed	Solution improved by deactivating satellites	- *	10**	3***	19	19
Fixed	Fixed	Minor or no-interventions	128	120	131	109	108

(*) *Using Bernese no attempts were made to improve the initial solutions.*
 (***) *For six flagged baselines it was not possible to obtain non-flagged solution.*

332

333 The results of the different software programs agreed quite well in the case of BMs that offer
 334 good observation conditions. On the contrary, for baselines involving BMs with unfavorable
 335 signal reception, significant differences resulted between the solutions of the different
 336 programs. For this reason, the comparison of the different software packages was done
 337 distinguishing between two classes of baselines: typical and problematic baselines. Two
 338 criteria were used for the classification of the baselines. The first criterion was the statistics of
 339 the solutions, i.e. the a-posteriori reference variance, the RMS and the standard error of the
 340 baseline components. Additionally, for the Bernese solutions we used as an additional

341 criterion the percentage of resolved ambiguities, which is being reported. The second
342 criterion for the classification of the baselines was the closure error. Instead of using loop
343 closures we computed the closure error based on the difference between the coordinates
344 resulted for each BM from each one of the two baselines available for that BM (from the two
345 nearest HEPOS stations). For our dataset this approach of computing the closure errors is
346 considered to yield more realistic results compared to loop closures, for two main reasons.
347 First, the two baselines used for each closure check are uncorrelated. If we had solved also
348 the baseline between the two HEPOS stations, each triangle would consist of three correlated
349 vectors. As known, three receivers measuring in parallel produce only two stochastically
350 uncorrelated baselines (Hofmann-Wellenhof et. al, 2008). Secondly, the baseline between the
351 HEPOS stations was in some cases twice as long as the baselines to the BMs, e.g. 69 km
352 between stations 043A and 007A (Figure 6). Baselines of such length cannot be precisely
353 estimated from occupations of one-hour duration. This would lead to increased loop closure
354 errors. In our study the horizontal closure error (dS) is:

$$355 \quad dS = \sqrt{(E_{RS_1} - E_{RS_2})^2 + (N_{RS_1} - N_{RS_2})^2} \quad (16)$$

356 where the indexes RS_1 and RS_2 denote the coordinates obtained from the baselines from
357 the nearest and the next nearest HEPOS reference station (RS), respectively. For the vertical
358 closure we used the absolute value of the difference between the ellipsoidal heights obtained
359 from each pair of baselines, i.e.:

$$360 \quad |dh| = |h_{RS_1} - h_{RS_2}| \quad (17)$$

361 Using the aforementioned criteria 39% of the baselines in Attica (11 among 28 baselines)
362 and 23% of the baselines in Thessaloniki (9 among 39 baselines) have been designated as
363 problematic. Tables 4 and 5 give the mean and maximum values of the horizontal and vertical
364 closure errors for the typical and problematic baselines in Attica and Thessaloniki,
365 respectively. Float solutions have been excluded from the computation of the results. The

366 mean values are depicted graphically in Figure 6 (Attica) and Figure 7 (Thessaloniki). In
 367 Addition, these figures include statistics computed over the entire sample of baselines for
 368 each area. The reason for this is that the problematic baselines were not common among the
 369 different software packages. Thus, the only way to directly compare the results is to examine
 370 the statistics over the same sample (all baselines). Comparing the results for the two areas it
 371 becomes obvious that the baselines in Thessaloniki offer slightly lower accuracy, on the order
 372 of few mm to 1 cm, compared to that of Attica. This is why we present our results
 373 distinguishing between the two areas. The lower performance in Thessaloniki can be mainly
 374 attributed to the fact that the measurements have been conducted with receivers of older
 375 technology (Leica SR520) compared to the receivers used in Attica (Topcon Hiper Pro).
 376

377 **Table 4:** Statistics of the horizontal closure error (values in m).

	Attica				Thessaloniki			
	Typical baselines		Problematic baselines		Typical baselines		Problematic baselines	
	Mean	Max	Mean	Max	Mean	Max	Mean	Max
BERN	0.028	0.060	0.075	0.120	0.054	0.152	0.108	0.205
GGO	0.012	0.031	0.015	0.045	0.037	0.117	0.045	0.125
GN	0.024	0.072	0.059	0.132	0.038	0.200	0.118	0.257
TBC	0.010	0.019	0.021	0.043	0.020	0.056	0.029	0.070
TT	0.012	0.027	0.034	0.104	0.020	0.050	0.062	0.140

378

379 **Table 5:** Statistics of the vertical closure error (values in m).

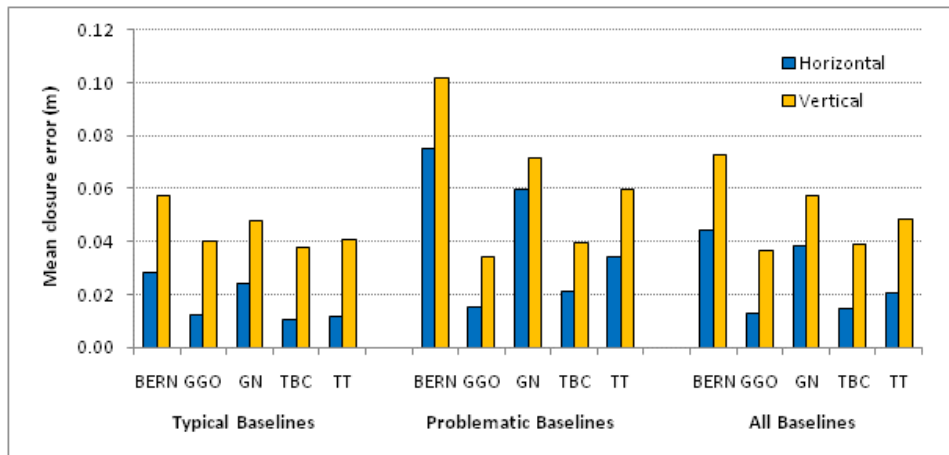
	Attica		Thessaloniki	
	Typical	Problematic	Typical	Problematic

	baselines		baselines		baselines		baselines	
	Mean	Max	Mean	Max	Mean	Max	Mean	Max
BERN	0.057	0.184	0.101	0.204	0.062	0.169	0.149	0.385
GGO	0.040	0.074	0.034	0.074	0.033	0.115	0.039	0.112
GN	0.048	0.202	0.072	0.185	0.052	0.162	0.101	0.239
TBC	0.038	0.085	0.040	0.075	0.040	0.098	0.049	0.155
TT	0.041	0.094	0.060	0.168	0.041	0.122	0.066	0.146

380

381

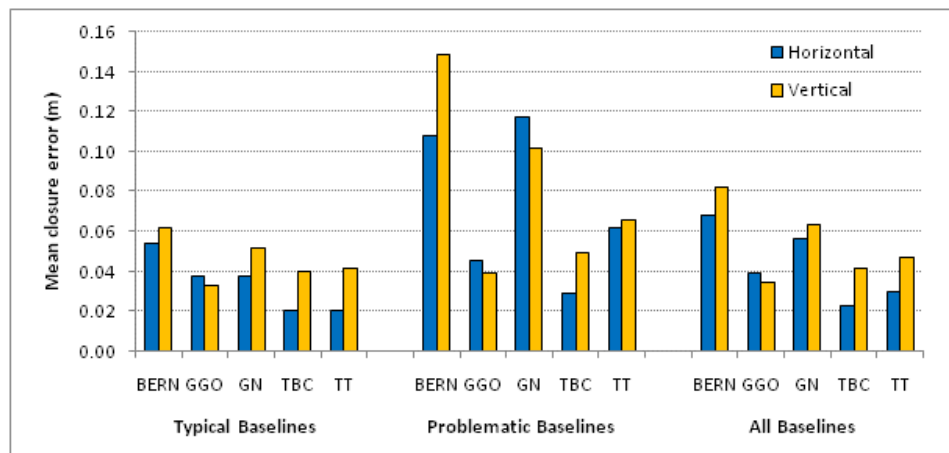
382



383

384 **Figure 6:** Mean horizontal and vertical closure errors for the baselines in the area of Attica.

385



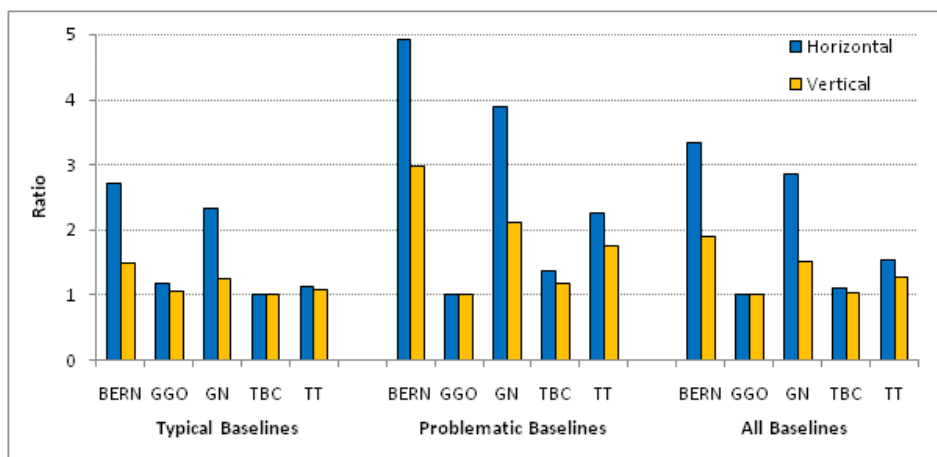
386

387 **Figure 7:** Mean horizontal and vertical closure errors for the baselines in the area of
 388 Thessaloniki.

389

390 In order to allow for a comparison of the closure errors of the different software we computed
 391 the ratio of the mean closure error of each software to the respective error of the best
 392 performing software. This comparison has been done separately for each group of baselines
 393 (*typical, problematic* and *all*) as well as for the horizontal and the vertical error. Figures 8 and
 394 9 give the computed ratios for the baselines in Attica and Thessaloniki, respectively. The best
 395 performing software can be easily recognized as its ratio equals 1. In each figure six bars are
 396 pointing at 1: for each group of baselines, one bar for the horizontal and one for the vertical
 397 error. In the case of Figure 9 within each group of baselines the lowest horizontal and vertical
 398 errors were obtained from the same software (TBC for the typical, GGO for the problematic
 399 and GGO for all baselines). On the contrary, in the case of more noisy observations in
 400 Thessaloniki (Figure 9) the best performance in horizontal and vertical closures within each
 401 group of baselines was achieved by different software.

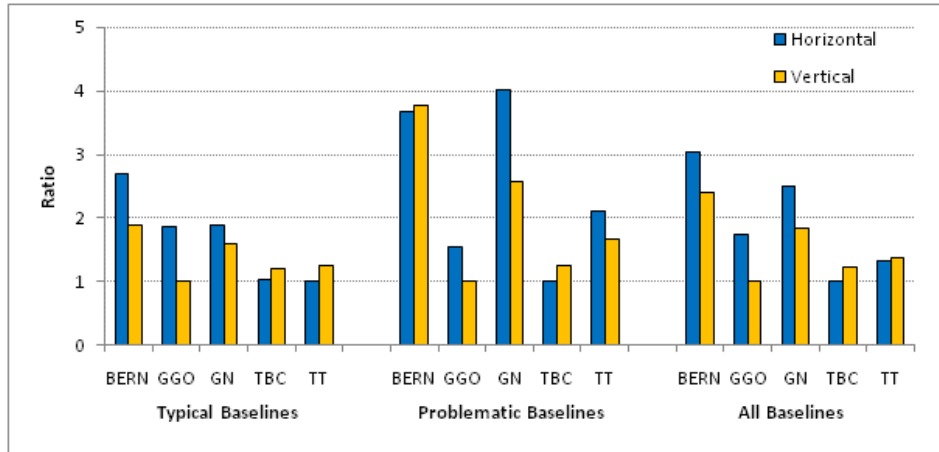
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403

404 **Figure 8:** Ratio of mean closure error of each software w.r.t. the best performing software,
 405 computed separately for dS, |dh|, typical, problematic and all baselines in Attica.

406



407

408 **Figure 9:** Ratio of mean closure error of each software w.r.t. the best performing software,
 409 computed separately for dS, |dh|, typical, problematic and all baselines in Thessaloniki.

410

411 4.3. Discussion of GPS results

412 Before discussing the results of the used software packages, we would like to stress that our
 413 purpose was not the assessment of the relative performance of the various software programs.
 414 Such comparisons require a much larger data set of baselines and the use of the latest
 415 versions of all programs, which was not the case in our study (see also Table 2 for the release
 416 year of each software). Actually, our goal was to demonstrate the importance of the
 417 processing software in the accuracy of the results, especially in the case of problematic
 418 baselines.

419 Examining the overall relative performance of the five software (Figures 6-7: columns *all*
 420 *baselines*) it becomes obvious that the four commercial software packages yield better results
 421 compared to Bernese. Of course, this conclusion does not reduce the worth of this well-
 422 acknowledged software, which undoubtedly belongs to the best scientific GNSS processing
 423 software worldwide. One should keep in mind that Bernese mainly focuses on the processing
 424 of measurements of long duration (e.g. daily occupations) collected at sites offering good
 425 observation conditions (e.g. reference stations) over long distances (baseline length of the
 426 order of several hundreds or thousands of kilometers). The detailed modeling of many errors

427 sources (ocean, atmospheric and solid earth tidal displacements, earth orientation variations,
428 satellite phase center offsets and patterns etc.) (Dach et al., 2007) is necessary for long
429 baselines, but does not actually improve the solution of short baselines, as these errors cancel-
430 out when forming double-differences. In addition, the long duration of the observations is
431 important for Bernese in order to perform realistic estimations, e.g. for the tropospheric
432 delay. On the other hand, commercial software packages are designed to process not only
433 data of good quality, but also problematic measurements collected under unfavorable field
434 conditions.

435 Among the four commercial programs GrafNet yielded more noisy results. GrafNet is part of
436 NovAtel's GNSS post-processing software package, which is well-acknowledged for
437 GrafNav, a kinematic baseline and Precise Point Positioning (PPP) processor based on a
438 Kalman filter. GrafNav and GrafNet use the same GNSS processing engine. This processing
439 engine is proven to provide great results for kinematic measurements (Dao et al., 2004; Bláha
440 et al., 2011). Examining Figures 6-7 (columns *typical* and *problematic baselines*) one may
441 conclude that for static observations of good quality, GN yields somewhat worse results
442 compared to the other commercial software packages, but in the case of problematic baselines
443 the results were up to 4 times worse. This could be attributed to the processing engine, which
444 is by design more suitable for kinematic measurements. Examining Figures 8-9 one may
445 conclude that the performance of the three other programs is roughly on the same level. For
446 example, GGO shows slightly better performance in the case of problematic baselines in
447 Attica. On the other hand, one baseline in Attica could not be solved by GGO, a fact that is
448 not reflected in the figure. In the case of problematic baselines in Thessaloniki TBC performs
449 significantly better than GGO. Observing the columns *all baselines* in Figures 8-9, we can
450 conclude that GGO, TBC and TT provide more-or less comparable results. If we consider
451 jointly the results in both areas, TBC shows the best performance. We attribute this

452 superiority mainly to the fact that TBC is the only software among the three programs that
453 gives detailed baseline processing report that contains a graphical representation of the
454 observation residuals. This functionality allows the detection and exclusion of noisy
455 observations, which considerably improves the solution. According to its manual TT has the
456 same capability, but it is available only if the “Advanced Module” for processing has been
457 licensed (Topcon, 2009). Regarding GGO, one could expect that the graphical representation
458 of residuals would be supported as this program is practically the same as LGO (Leica Geo
459 Office). However, comparing the two software packages one can see that certain
460 functionalities of LGO are not available in GGO (Leica, 2010).

461 Examining the closure errors of typical and problematic baselines (Figure 6) we can see that -
462 although the horizontal errors are lower than the vertical- they increase up to 3 times in the
463 case of problematic baselines (TT). On the contrary, the increase of the vertical errors is
464 limited to a factor of 1.7 (BERN). For the sake of clarity, we would like to stress that in the
465 case of problematic baselines the vertical errors are still larger. However, the accuracy
466 degradation caused by the problematic observations is higher for the horizontal component.
467 This is a result of practical importance for the professional surveyors, who often measure in
468 difficult environments and they are mainly interested in the horizontal accuracy.

469 Figure 6 verifies the general rule, which states that the vertical accuracy of GPS baselines is
470 considerably lower than the horizontal accuracy. Looking at Figure 7 we can find some
471 exceptions to this rule. More specifically, GGO provided smaller vertical errors for all group
472 of baselines in Thessaloniki and GN showed similar behavior in the case of problematic
473 baselines. To some extent these results could be explained by the fact that the observations in
474 Thessaloniki are characterized by increased noise, as discussed above. As explained in the
475 previous paragraph, the relationship between horizontal and vertical precision alters in the
476 case of problematic observations. However, even in the case of problematic observations the

477 vertical errors still remain generally higher. Thus, the different behavior of GGO and GN is
478 believed to originate from the particular processing algorithms implemented in each software
479 package. This investigation requires detailed comparison of the different GNSS processing
480 engines, a task that is beyond the scopes of this paper.

481

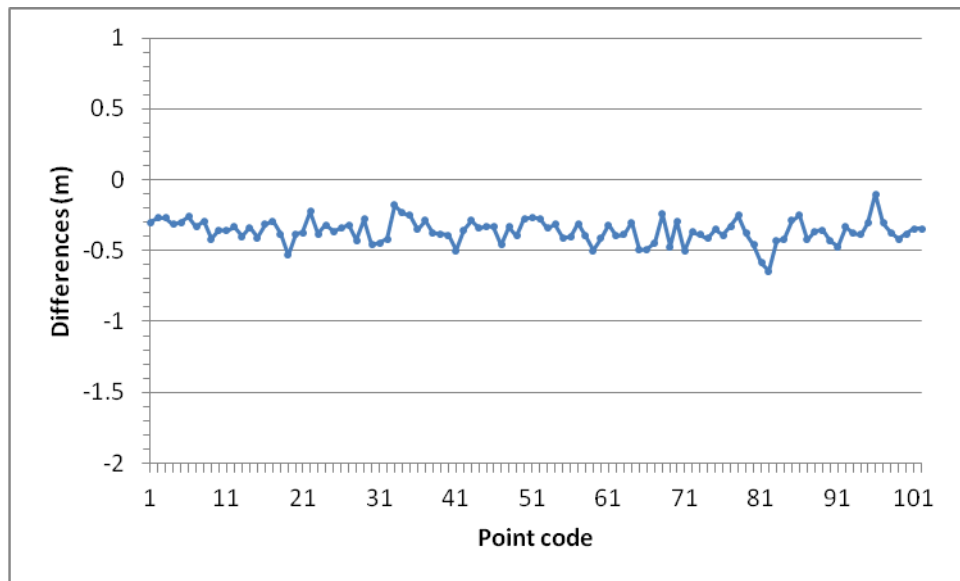
482 **4.4. Comparisons with global geopotential model EGM2008**

483 The initial stage of the validation at GPS/leveling benchmarks is based on comparisons with
484 external information. GPS/leveling provides the geometric connection between different
485 height systems (geometric/ellipsoidal and orthometric height). According to eq. (14) a
486 geometric estimation of the geoid can be derived using ellipsoidal and orthometric height
487 information. The determination of this “geometric” geoid is directly comparable to the
488 “physical” one derived from a geopotential model of high accuracy and resolution. As
489 presented in the theoretical part, the resolution of the geopotential model is based on the
490 degree and order of its coefficients expansion and its accuracy on the commission and
491 omission errors estimated during the adjustment process (Hofmann-Wellenhof and Moritz,
492 2005). It should be kept in mind though that such a “geometric” geoid model is of limited, if
493 any, theoretical rigorousness. This is due to the fact that the formed $h - H$ differences do not
494 realize the geoid, i.e., a physical surface of constant gravity potential (W_0). They rather
495 realize the difference between the two heights along the vertical lines with any systematic
496 distortions due to the different datum of h and H . The major problem of the established Greek
497 vertical datum is its systematic distortion due to the largely unknown accuracy of the BM
498 orthometric heights. The non sufficient documentation on the adjustment procedure
499 (constraints type and number) and the lack of the covariance matrix estimation

500 The global geopotential model used in the comparisons is the state-of-the-art spherical
501 harmonics expansion geoid model based on various data sources combined, Earth

502 Gravitational Model 2008 - EGM2008 (Pavlis et al., 2012). This model incorporates
503 optimally surface gravity observations, satellite altimetry data and newly available products
504 from gravity dedicated satellite missions (GRACE). The spherical harmonic expansion of
505 EGM2008 reaches degree 2190 and order 2159, resulting in a spatial resolution of 5 arc
506 minutes. In the present study, EGM2008 contribution is utilized up to degree and order of
507 expansion 2159. According to recent studies, the maximum degree 2190 showed only minor
508 improvements in the Hellenic area (Tziavos et al., 2010). Figure 10 presents the differences at
509 the 103 benchmarks in Attica region. A mean value of -0.362 m is calculated. This bias
510 represents the W_0 offset of the Greek vertical datum with respect to EGM2008. The internal
511 accuracy of the procedure can be expressed by the standard deviation of the differences
512 computed ± 0.083 m in Attica region.

513



514

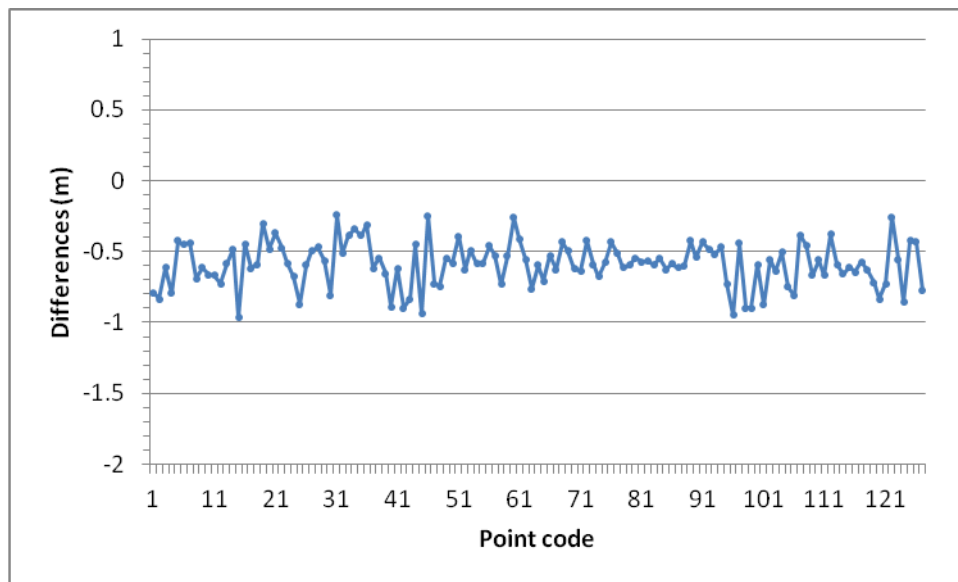
515 **Figure 10:** The differences between GPS/levelling and geopotential model derived geoid at
516 Attica test area after blunders removal

517

518 Approximately the same situation is presented in the test area of Thessaloniki. The
519 differences between GPS/leveling and the GGM geoid heights are charted in Figure 11. The

520 statistics of the 127 point differences demonstrated a mean of -0.588 m and ± 0.158 m
521 standard deviation. The clarification of a bias difference (approximately 0.20 m) between
522 Attica and Thessaloniki area results is part of our future research plan related to the
523 unification of the Greek Local Vertical Datum (LVD). At first glance, it can be attributed to
524 datum inconsistencies in the vertical datum. The standard deviation of the differences ± 0.158
525 m reveals an accuracy degradation from the results of Attica which is attributed to the fact that
526 the study area of Thessaloniki has rougher terrain and it is characterized by higher elevations.
527 Hence, orthometric heights are, naturally, of lower accuracy.

528



529

530 **Figure 11:** The differences between GPS/levelling and geopotential model derived geoid at
531 Thessaloniki test area

532

533 **5. Conclusions – Future Plans**

534 The investigation of the internal as well as the external accuracy of the Hellenic vertical
535 network is the main goal of E.LE.V.A.T.I.O.N. project. Two test areas are chosen and the
536 initial assessment of the internal accuracy of the network is based on GPS measurements at
537 benchmarks with known orthometric heights. Different GPS processing software packages

538 are used and compared to each other. The global geopotential model EGM2008 is utilized for
539 the assessment of the external accuracy of the network. Two test areas are chosen in Central
540 and Northern Greece containing 230 benchmarks in total.

541 Based on the discussion of the GPS processing, some conclusions related to the performance
542 of different software packages can be drawn. In the case of the baselines tested here (short
543 baselines, a few tens of kilometers in length, observed for 1 hour) the commercial software
544 packages perform better than the scientific one. The requirement of increased amount of data
545 for the proper modeling of a large number of parameters estimated by the scientific software
546 is the main reason for its reduced performance. Under unfavorable measurement conditions
547 (reduced satellite visibility and/or poor signal reception) there are noticeable differences in
548 the performance of the various software packages. Differences exist among the commercial
549 software packages based on the solution strategy of each one of them, depending on the
550 baseline length and the observation period. Some of these differences can be attributed to the
551 processing engine, which is by design more suitable for kinematic measurements than for
552 static ones.

553 The difficult measurement environment clearly affects the precision of the final result. This
554 fact stands for all software packages used in our study. The precision degradation is found
555 higher for the horizontal coordinates rather than for the heights, as the vertical component is
556 always estimated with reduced accuracy. This fact underlines the importance of the
557 observation conditions during a GPS campaign. A careful planning of the measurements is of
558 great importance for high precision applications. Nevertheless, generally speaking, the
559 horizontal closure errors are smaller than the vertical closures. However, certain software
560 programs provided slightly better results in the vertical component. This remark requires
561 further investigation.

562 The validation of the vertical datum in both test areas is performed using external information
563 from the state-of-the-art global geopotential model EGM2008. The results in Attica show an
564 agreement between “geometric” and “physical” geoid of 8.3.cm, in terms of the standard
565 deviation of the differences. In Thessaloniki, this agreement is 15.8 cm. A bias between the
566 average difference of Attica and Thessaloniki is observed, which can be attributed to the
567 datum offset between the Greek datum and EGM2008. This bias presents different
568 characteristics in Attica than in Thessaloniki, resulting a 20 cm offset, approximately,
569 between the average differences at the two areas. The abovementioned offset is related to the
570 LVD used in each area and it is the subject of ongoing work. It should be noted that due to
571 the absence of sufficient documentation and the repeated partial adjustments performed since
572 its creation, the actual accuracy of the Hellenic vertical datum is largely unknown. The use of
573 additional geopotential models, especially the recent available models from GOCE satellite,
574 will contribute to the efficient validation of the height datum with respect to its spectral
575 characteristics.

576

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583

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